

Please see our replies to the three reviewers in the related three single files. We believe we have fully replied to their requests where pertinent. In some cases we have pointed out why their suggestions (e.g. the triple co-location technique) were not pertinent.

The authors

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Accuracy of altimeter data in inner and coastal seas

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35

Abstract

36 We carry out an inter comparison among four different altimeters, Cryosat, Jason2, Jason3,
37 Sentinel-3. This is done checking the altimeter data versus the same wind and wave model
38 results of a given area, the Mediterranean Sea, for one year period. The four datasets are
39 consistent for wind speed, but they show substantial differences for wave heights. The
40 verification of a Sentinel-3 pass close to coast in the Northern Adriatic Sea shows irregular spiky
41 large wave height values close to coast. The problem worsens using high frequency altimeter
42 data.

43

44 1 – Altimeter data

45 There is no doubt that satellite radar altimetry has revolutionized oceanography with the
46 continuous and abundant flow of data during the last three decades or so. The related surface
47 wind speed and significant wave height data have provided both a crucial information for data
48 assimilation in, and validation of, model activity and results (see, among others, Abdalla, 2007,
49 2016) and substantial and prolonged information to be used for global statistics over the oceans.

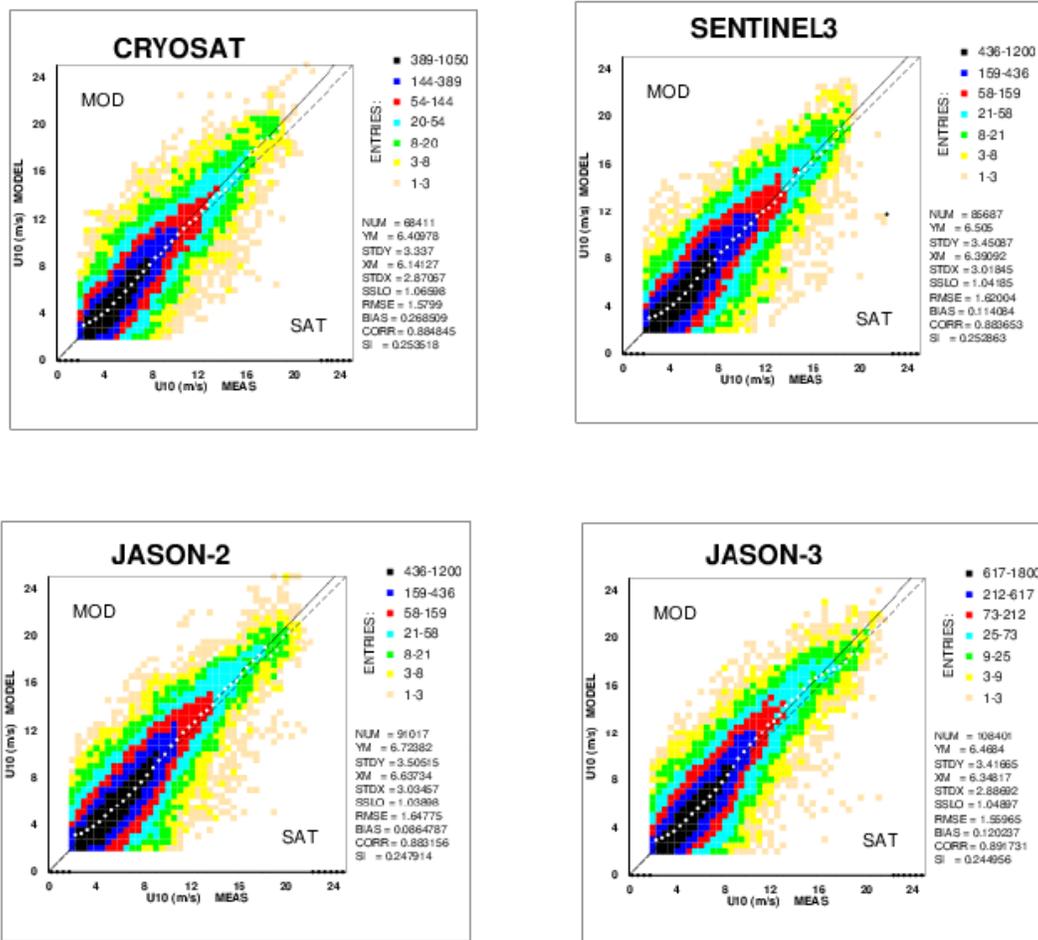
50 Although officially calibrated, a careful intercompariosn strongly suggests the data from
51 different instruments require specific attention and calibration. See in this respect the keen and
52 prolonged analysis by Young et al. (2017). Starting in 1986 with Geosat, the real continuous
53 flow began in 1991 with the launch of ERS-1, followed in time by Jason1, ERS-2, Jason2,
54 Jason3, Envisat, Altika, Cryosat, ending (for the time being) with Sentinal-3. See also Passaro et
55 al (2014).

56 The two different principles of interaction with the sea surface for wind and wave information
57 retrieval (respectively back-scattering and specular reflection) imply different calibrations for the
58 two signals, calibrations sensitive to the average conditions where the operation has been done.
59 See in this respect the valuable work by Queffeulou and Bertamy (2007). Being this mainly
60 versus buoy data in the oceans, it is correct to wonder if this calibrations hold also in the rather
61 different conditions of the inner seas. This is particularly true in view of the use of Sentinel-3 in
62 very coastal waters in the attempt to push the use of altimeter data till very close to coast,
63 certainly much closer than the 20-30 km distance of the classical altimetry.

64 In this short paper we make an intercomparison between the wind and wave data from four
65 different altimeters and the results of two high resolution meteorological and wave models. The
66 area is the Mediterranean Sea. We use Cryosat, Jason2, Jason3, Sentinel-3 data (henceforth Cy,
67 J2, J3, S3). The period is the twelve months from July 2016 till June 2017. The models are
68 COSMO for meteorology (see www.cosmo.model.org/content/model/default.htm) and WAM for
69 waves (see the historical Komen et al., 1994, and the more updated Janssen et al., 2005). The
70 related operational system is the combined effort of the Italian Meteorological Service and the
71 Institute of Marine Sciences (ISMAR-CNR). COSMO is run at 7 km, WAM at 0.05° resolution.
72 The system provides twice daily three-day forecasts at hourly interval. A full description of the
73 system and its accuracy is available at Bertotti et al. (2013). For the altimeter-model inter
74 comparison we have used the first 12 hour forecasts of the twice-a-day operational activity
75 (hourly fields). The model data have been, bi-linearly in space and linearly in time, interpolated
76 at the position and time of each altimeter datum. The Cy, J2, J3 data have been retrieved from
77 the Delft University website <http://rads.tudelft.nl/rads/rads.shtml>. The S3 data from
78 <https://coda.eumetsat.int>.

79 We stress that for the most part, rather than on the comparison with the model data, the analysis
80 is based on an an, although indirect, intercomparison among the different altimeters. We present our

81 analysis and results in the next Section 2. In Section 3 we focus on an example of use of S3 data
 82 very close to coast. We summarize our conclusions in the final Section 4.



83
 84 Figure 1 – Scatter diagrams of the COSMO model wind speeds vs the Cy, S3, J2, J3 altimeter
 85 data. The area is the Mediterranean Sea. The continuous lines show the respective best-fit slopes.
 86 Dash lines would be the perfect fit.

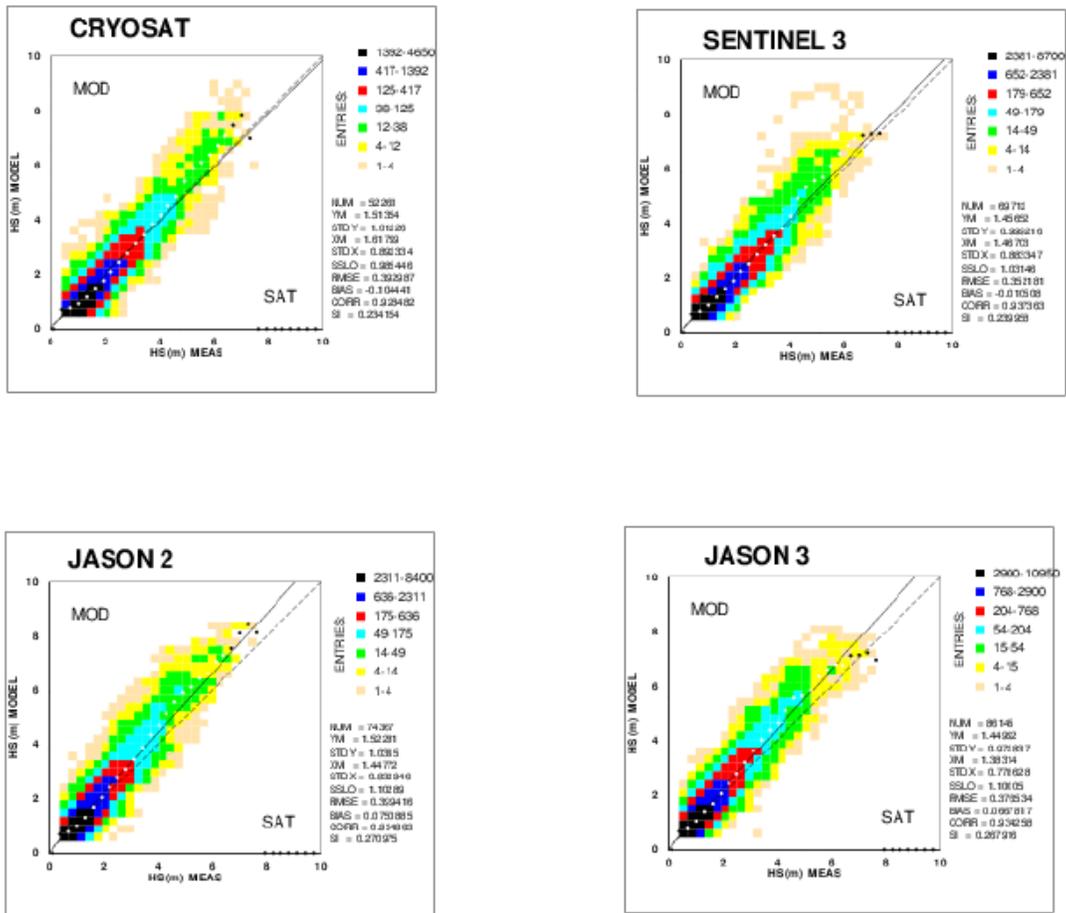
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88 2 – How much altimeters differ one from the other one

89 Figure 1 provides the scatter diagrams and related statistics for the model surface wind speed vs
 90 the altimeter data for the four considered satellites. As usual, the colours (scale on the right side
 91 of each panel) indicate the number of data in each pixel.

92 Varying between 4 and 6%, there is a clear relative overestimate by the model. There is a rather
 93 large scatter, with the scatter index SI typically at 0.25. There is some indication of a larger
 94 overestimate in the higher value range. However, for the present purpose we focus on the

95 altimeter data, and the results in Figure 1 suggest a consistent performance among the four
 96 altimeters.



97
 98 Figure 2 – As Figure 1, but for the significant wave height.

99 Things are rather different when we consider the significant wave height H_s . The related results
 100 are in Figure 2. There are obvious differences among the four altimeters, summarized for waves
 101 in Table 1.

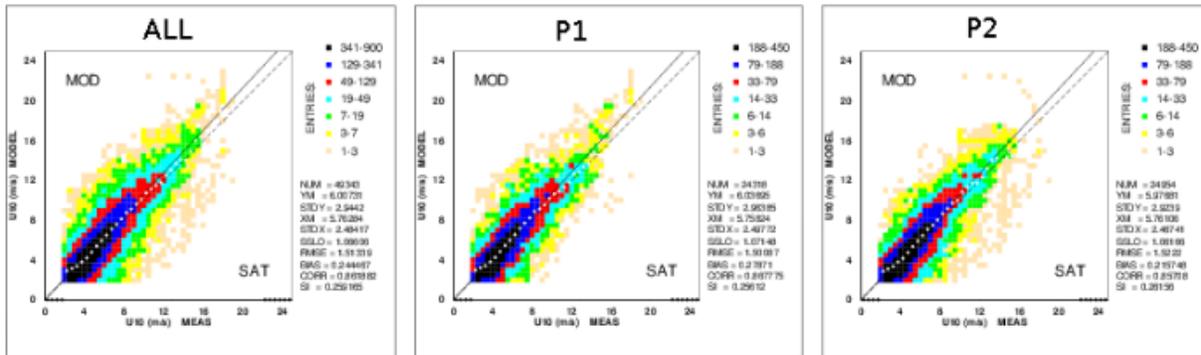
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alt	Cy	J2	J3	S3
sslo	0.98	1.10	1.10	1.03
SI	0.23	0.27	0.27	0.24

103

104 Table 1 – Symmetric best-fit slope (sslo) and scatter index (SI) of the wave model data versus
 105 altimeter ones. Cy = Cryosat, J2 = Jason2, J3 = Jason3, S3 = Sentinel-3.

106



107

108 Figure 3 – As Figure 1, but for only the Cryosat data. Left panel (ALL) full dataset, P1 and P2
 109 each one complementary half of the data selected by random sampling.

110 There is a 12% wave height best-fit difference between the two Jason altimeters and Cy, and still
 111 7% versus S3. Our arguing is the following. Given that the different altimeters have been
 112 compared with the same model data (area and period), these results must reflect differences
 113 among the four instruments. However, it can be argued that the four instruments have not
 114 measured exactly the same wave conditions (in space and time), each satellite sampling at
 115 different times and positions, hence on different wave conditions. To explore this possibility, we
 116 have split each altimeter dataset in two halves with a random sampling of the different passes,
 117 then evaluating a new statistics for each one of the two half of the data. As an example we show
 118 in Figure 3 the related results for Cy (similar results hold for each satellite). There is hardly any
 119 difference. More in general for each altimeter the differences among the full and half statistics
 120 are less than 2% of the single statistical figures. Therefore the results in Table 1 are fully
 121 representative of the situation.

122

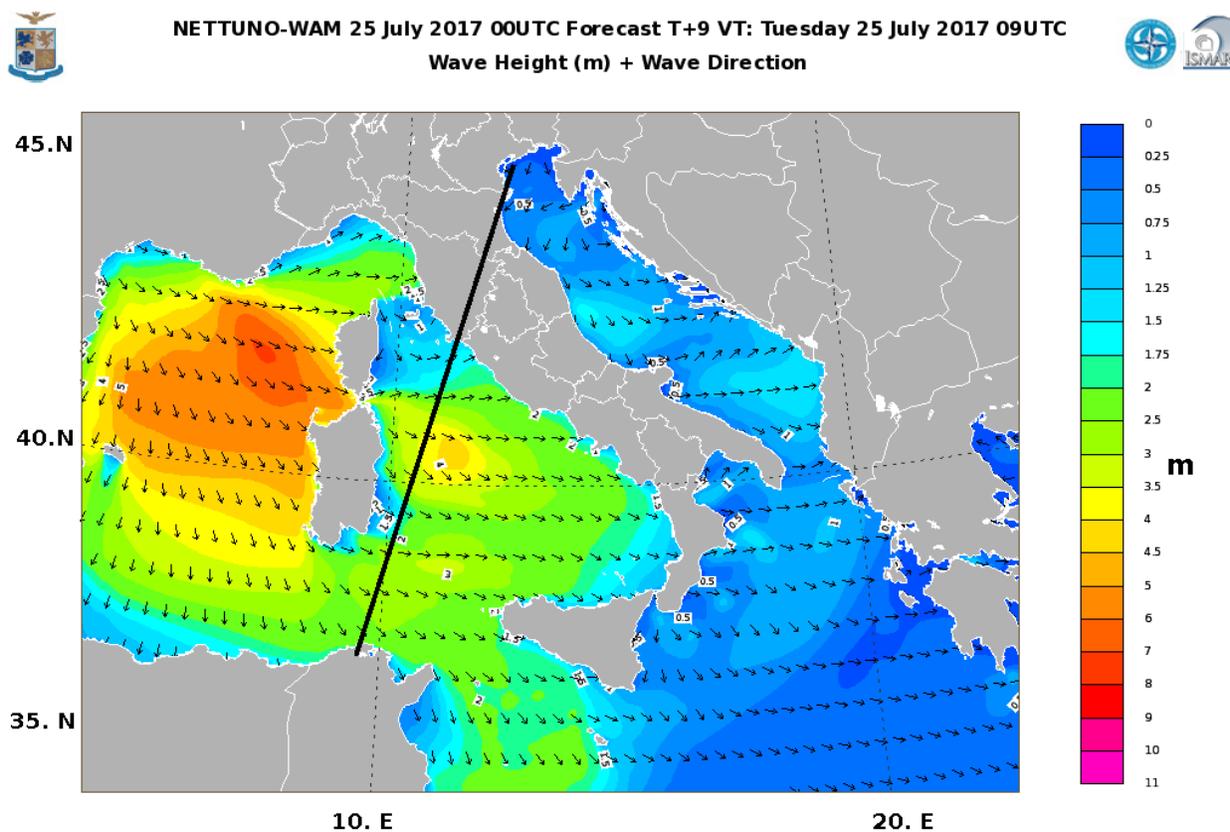
123 3 – A Sentinel-3 coastal track

124 The Sentinel-3 altimeter acquisition system, somewhat an evolution of the SIRAL mounted on
 125 Cryosat-2 satellite, is claimed to enhance the accuracy of sea state measurements close to the
 126 coast. S3 altimeter, SRAL, has two distinct operational modes: the conventional LRM and the
 127 high along-track resolution, or SAR mode (www.sentinel.esa.int). The latter is synthesized from
 128 a composition of 64 Ku-band pulses and two c-band ones. Operating in SAR mode, the SRAL
 129 along track resolution is of the order of 300 m with a large (up to 10 km) lateral swath. The
 130 reduced sampling area implies obviously a higher noise in the signal, but, especially when flying

131 perpendicularly to the coastline, it allows in principle to go much closer to it with meaningful
132 data. We explore this possibility analyzing one pass in the Northern Adriatic Sea.

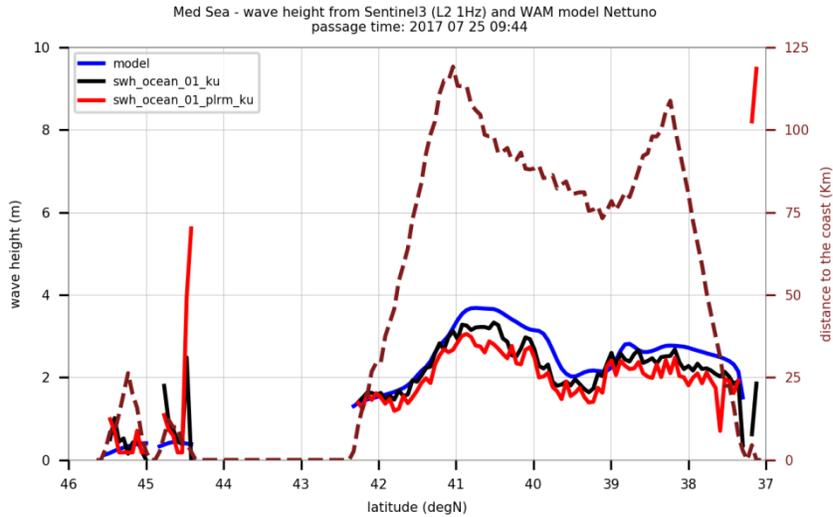
133 The proposed sample is a level 2 product provided from the Copernicus service with identifier
134 “S3A_SR_2_WAT____20170725T094431_20170725T094658_20170725T120008_0146_020_
135 193____MAR_O_NR_002.SEN”. The NetCDF file contains also PLRM (pseudo low res
136 mode) data, which is intended as a simulation, starting from Ku-band pulses, of the classical
137 altimeter sampling strategy, LRM.

138 **Figure 4 shows the** ground track during an S3 25 July 2017 descending pass over first the
139 Adriatic then the Tyrrhenian Sea. There was a severe mistral storm in the Western Mediterranean
140 (see the H_s scale on the right), but only a tiny bit of it passing between Corsica and Sardinia was
141 touched by the pass. The model and altimeter data are in Figure 5, latitude decreasing, hence
142 following the satellite, from left to right. For a short moment we focus on the Tyrrhenian results,



143

144 **Figure 4 – Ground track of a descending pass of Sentinel-3 altimeter. The background colours**
145 **and arrows show the significant wave height distribution.**

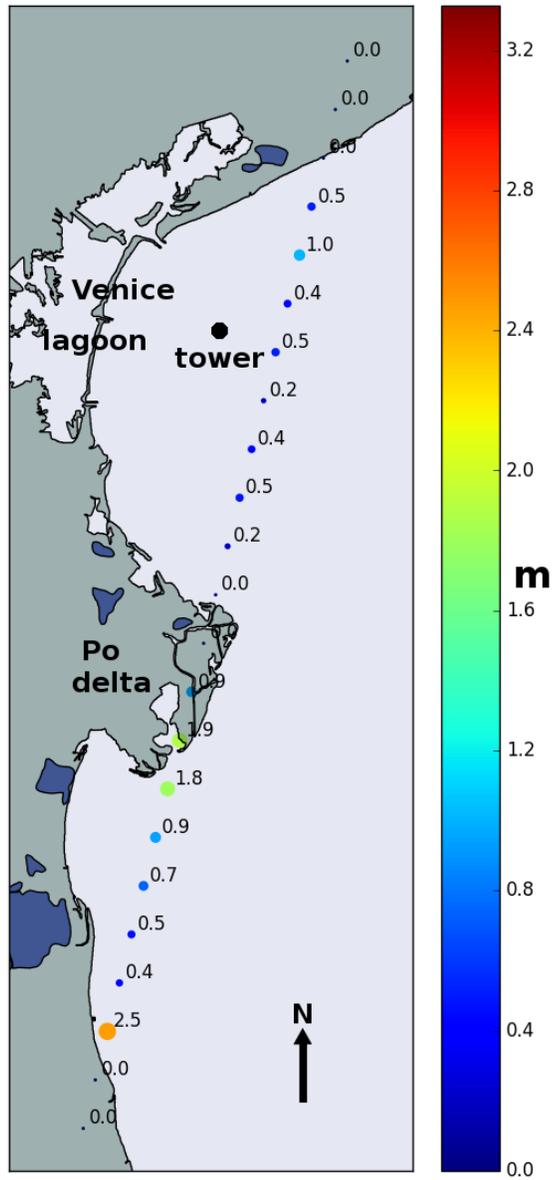


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147 Figure 5 – Intercomparison, along the ground track in Figure 4, between the model significant
 148 wave heights and the measurements (ku-, plrm-ku- corrected H_s) by the Sentinel-3 altimeter. The
 149 dash line shows the distance (km) of each measurement from the closest coast.

150

151 the model H_s following well the measured quantity. The dash line (right scale) shows the
 152 distance from the closest coast (km). Note the altimeter spikes when exiting and entering land. In
 153 this respect we zoom on the short passage on the Adriatic Sea (the first short section in Figure 5),
 154 passage geographically better represented in Figure 6. We recognize the Venice lagoon (about 50
 155 km long) and the protruding Po river delta intersected by the descending satellite ground track.
 156 Dots and close-by numbers represent the 1 sec S3 H_s data (ku-band) Note the incoherent data
 157 when passing on the Po delta and when entering land again shortly after. The oceanographic
 158 situation is in Figure 7. There is a very mild wind sea from North-East with significant wave
 159 height close to, mostly less than, 0.5 m (product of the operational ISMAR Adrioper wave
 160 forecast system, see Bertotti and Cavaleri, 2009). An independent validation (not shown) of the
 161 model results for this day is provided by the data regularly recorded at the ISMAR
 162 oceanographic tower (Cavaleri, 2000), located 15 km off the Venice lagoon (see Figure 6).
 163 The model-measurement H_s difference close to the time of the pass is less than 10% that, on the base
 164 of previous experience and validation (Bertotti and Cavaleri, 2009), we take as characteristic of
 165 the overall local field, hence of also the model data corresponding to the S3 ones.



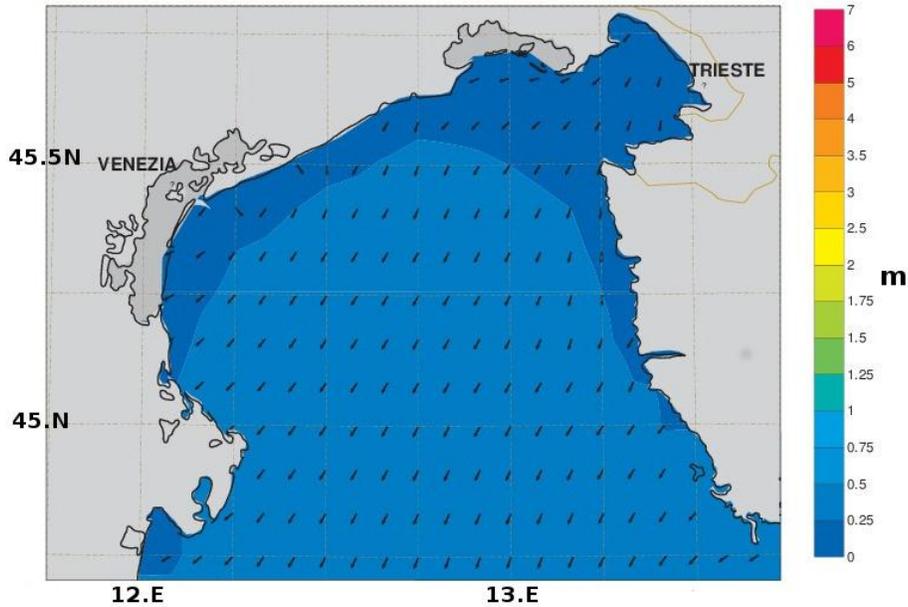
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167 Figure 6 – Detailed geometry, focused on the Adriatic Sea, of the area of the pass in Figure 4, in
 168 more details in Figure 7. The positions and the corresponding SAR mode (swh_ocean_01_ku)
 169 altimeter significant wave height values (m) are also shown.

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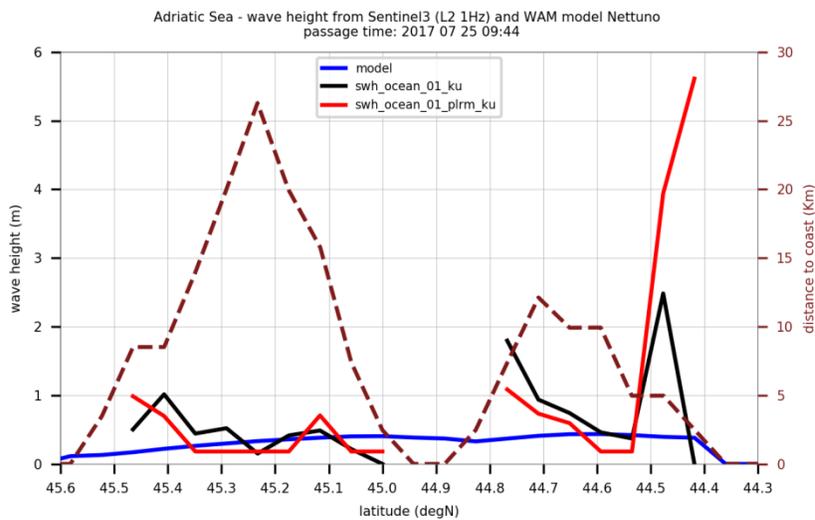


Previsione per il : 25 Luglio 2017 ore 09 UTC
 Altezza d'onda significativa (m) + direzione media



171

172 Figure 7 – Wave field (very mild conditions) in the Northern Adriatic Sea at 09 UTC 25 July
 173 2017. The arrows show the significant wave height and mean direction. The modelled maximum
 174 wave height in the field is close to 0.5 m. Wind and waves were from East-North-East.

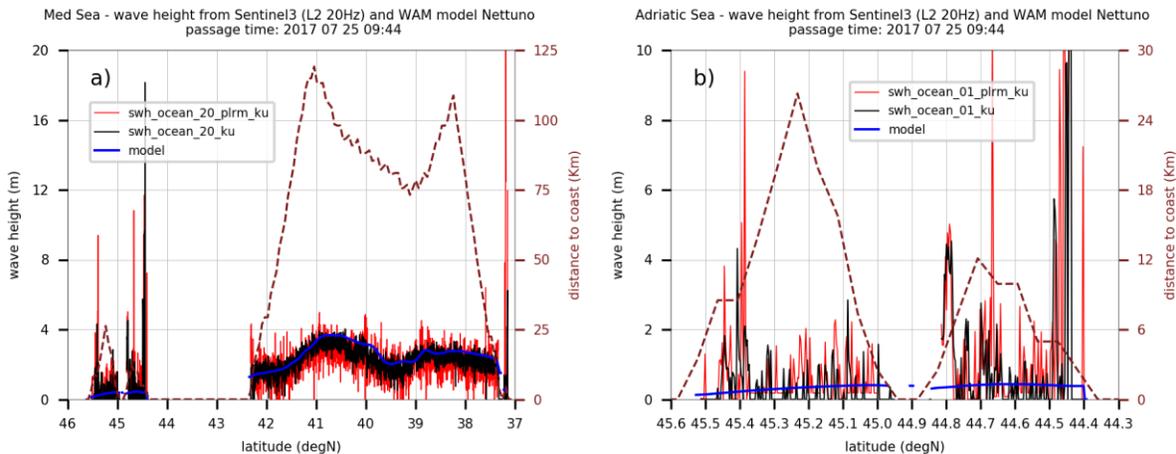


175

176 Figure 8 – Intercomparison, for the pass in Figure 6 and the time of Figure 7, between the
 177 Sentinel-3 ku-, plrm-ku- wave heights and the corresponding wave model results. The dash line
 178 shows the distance from the closest coast (km).

179

180 In Figure 8 we show the detailed comparison among the three different (ku-, plrm-ku-) bands
 181 and model H_s values plus the distance (km/10) from the closest coast. There are some obviously
 182 absurd values by S3 in the two signals, more in the plrm-ku, this when the distance from the
 183 coast was less than 10 km.



184

185 Figure 9 – As Figure 8, but for the 20 Hz altimeter data. Panel a is for the full pass, panel b for
 186 the Adriatic Sea section (see Figure 8 for comparison).

187 Finally in Figure 9 we explore the 20 Hz data. Panel a shows the whole pass (the corresponding
 188 of Figure 5). Panel b focuses on the Adriatic Sea. The noise of the signal is evident, also when
 189 the distance from the coast was about 20 km. There is a very large variability of the S3 altimeter
 190 signal also in the Tyrrhenian Sea, a variability that cannot be justified by geophysical reasons,
 191 and it is therefore natural to associate to the instrument and to the sampling variability. Again the
 192 S3 approach seems to lead to very large H_s values also when the distance from the coast
 193 approaches the classical 20-30 km limit of standard altimetry.

194

195 4 – Summary

196 Following the extensive availability of altimeter data for both wind and waves, and their
 197 relevance in validating model results, we have explored the consistency of the data from four
 198 different altimeters. Lacking the possibility of extensive triple-collocation analysis (the datasets
 199 would be too small for meaningful results), we have followed a different principle, i.e.: we take
 200 the model data as reference and, without arguing about right or wrong, we explore how each
 201 satellite fits the model data. Should the altimeters being consistent to each other, each altimeter
 202 vs model fit should provide the same best-fit slope. On a different line of action we have
 203 explored the value of S3 data close to coast. In summary, we have carried out two tests: 1) an
 204 extensive one on four different altimeters (Cy, J2, J3, S3, see Table 1), and 2) a sample one on
 205 one S3 pass. The purpose of 1) was an, indirect but significant, intercomparison among these

206 four altimeters. 2) was meant to explore one case of sampling by S3 in coastal waters. In this
207 case we have also checked the value of the 20 Hz data. We itemize our results as follows.

208 1 - the surface wind speed values derived from the four altimeters are consistent to each other,
209 differing on the average less than 2%,

210 2 - large differences are found in a similar intercomparison for the significant wave height H_s .
211 There is on average a 12% difference between the best-fit slopes of Cy and the J2-J3 data, the
212 latter ones measuring larger wave heights. The S3 values lay more or less in the middle,

213 3 - the S3 1 Hz data close to coast are noisy, with spikes of obviously wrong large values. Use of
214 20 Hz seems to increase the noise, wrong large values appearing also relatively far (20 km) from
215 the coast,

216 4 - the use of 20 Hz leads to a high variability of the H_s data also in the open sea, far from the
217 coasts, implying this variability is practically associated to the instrumental measurement and to
218 its sampling variability.

219

220 **Acknowledgements**

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224 paper.

225

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253 **Figure captions**

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